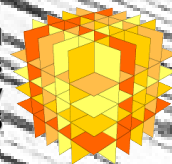


Single Phase Rocks

Mark Jessell

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Marie Curie Summer School
Knowledge Based Materials
Hürtgenwald, Germany
August 8-17th 2005



Compositional Structure of the Earth

thickness

composition

0-50 km

Crust

O, Si, Na, Ca, K, Fe, Mg, Al
Silicates, Carbonates, Oxides, Halides

2840 km

Mantle

O, Si, Fe, Mg Silicates & Oxides

2260 km

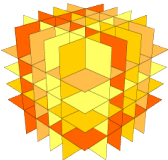
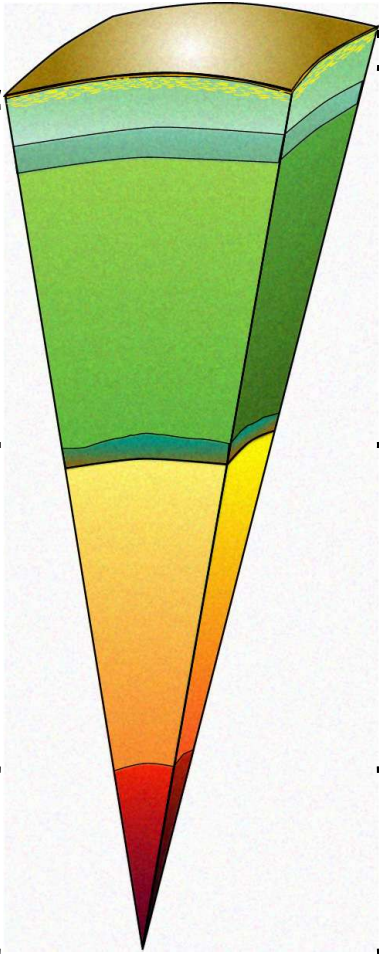
Outer core

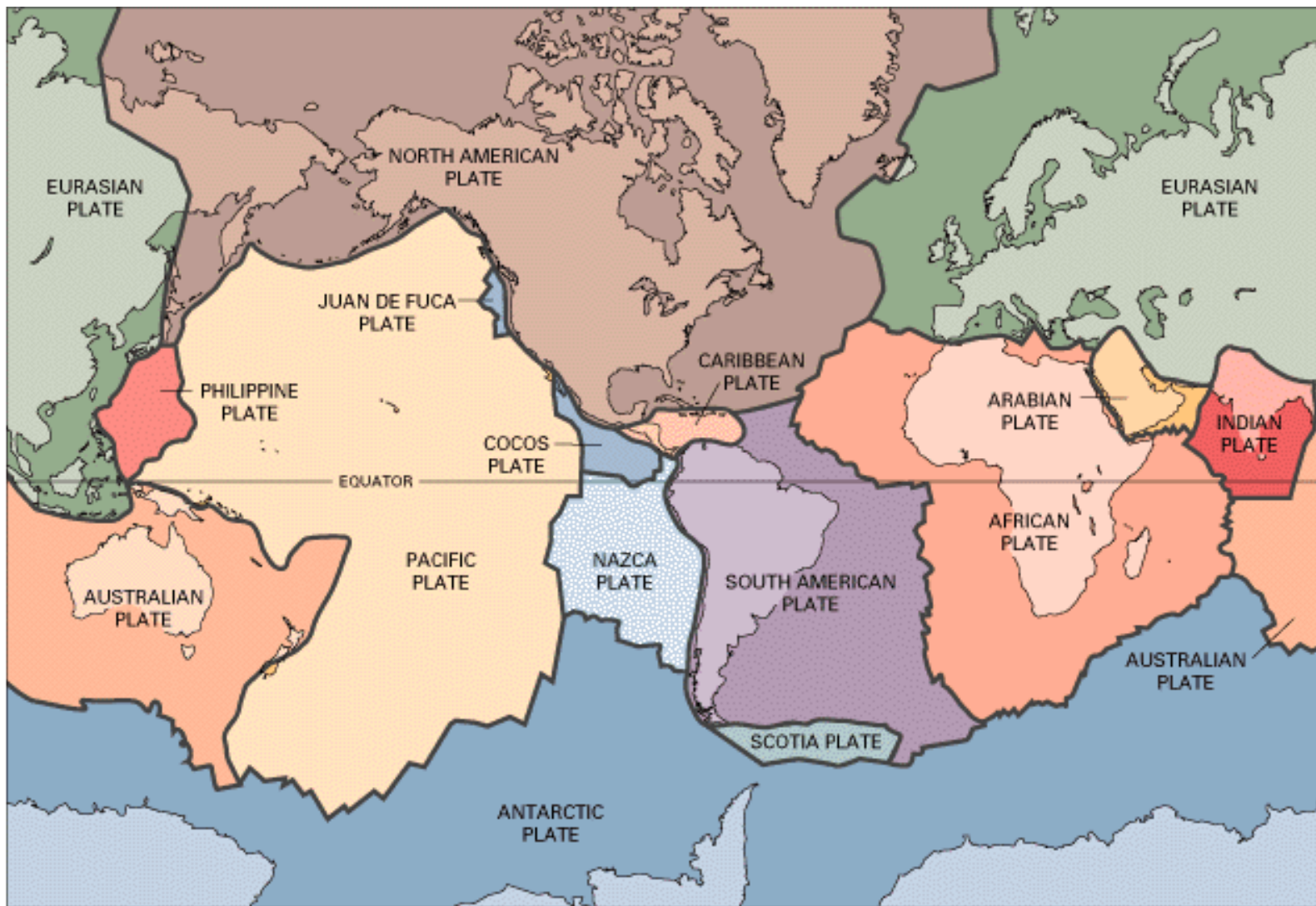
"Fe₂O", Sulfur, Nickel

1220 km

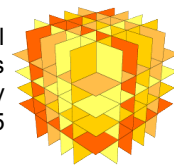
Inner core

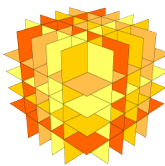
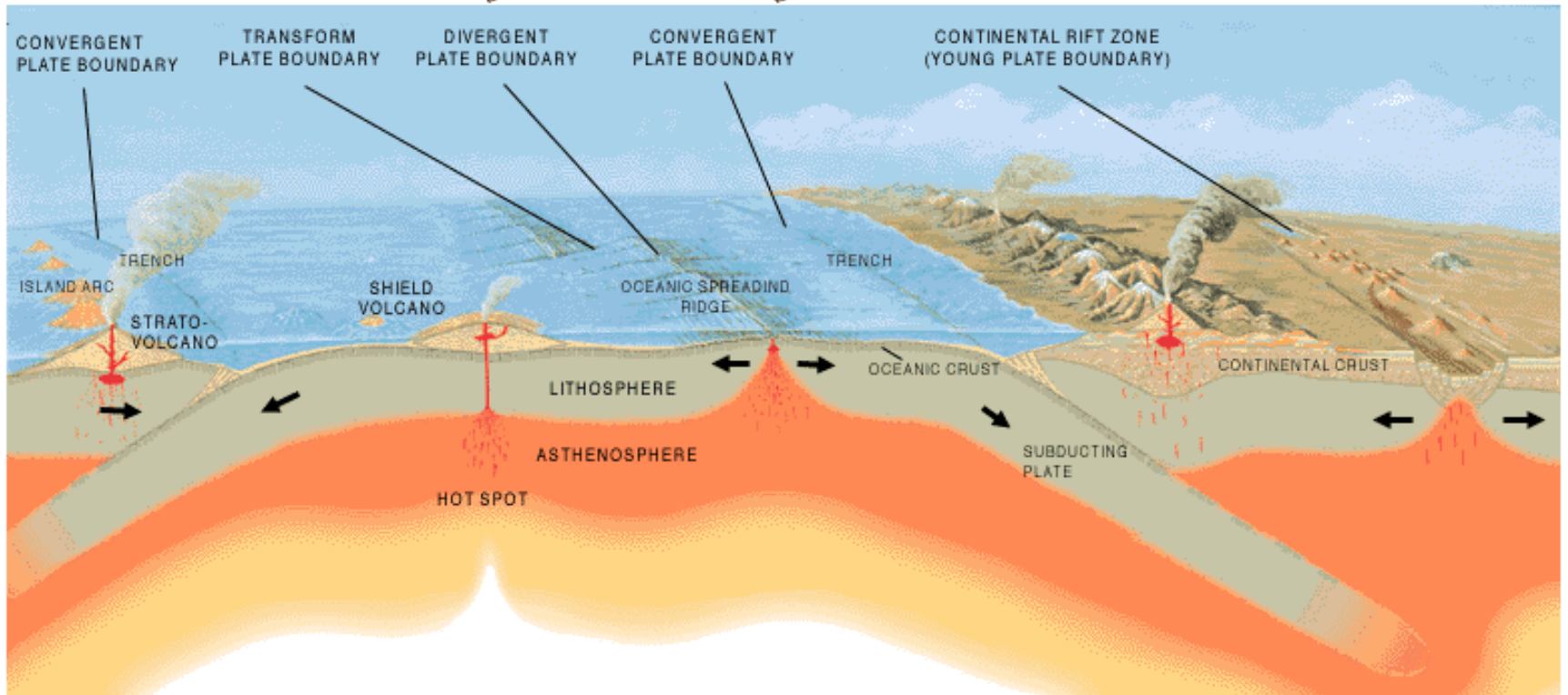
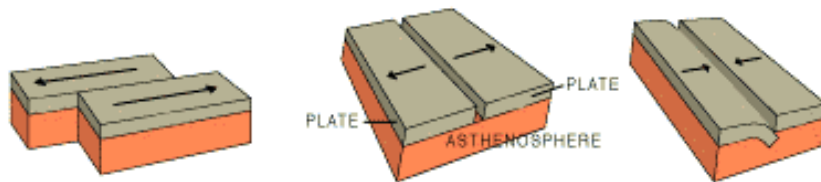
Pure Fe-Ni alloy, perhaps 6% Ni



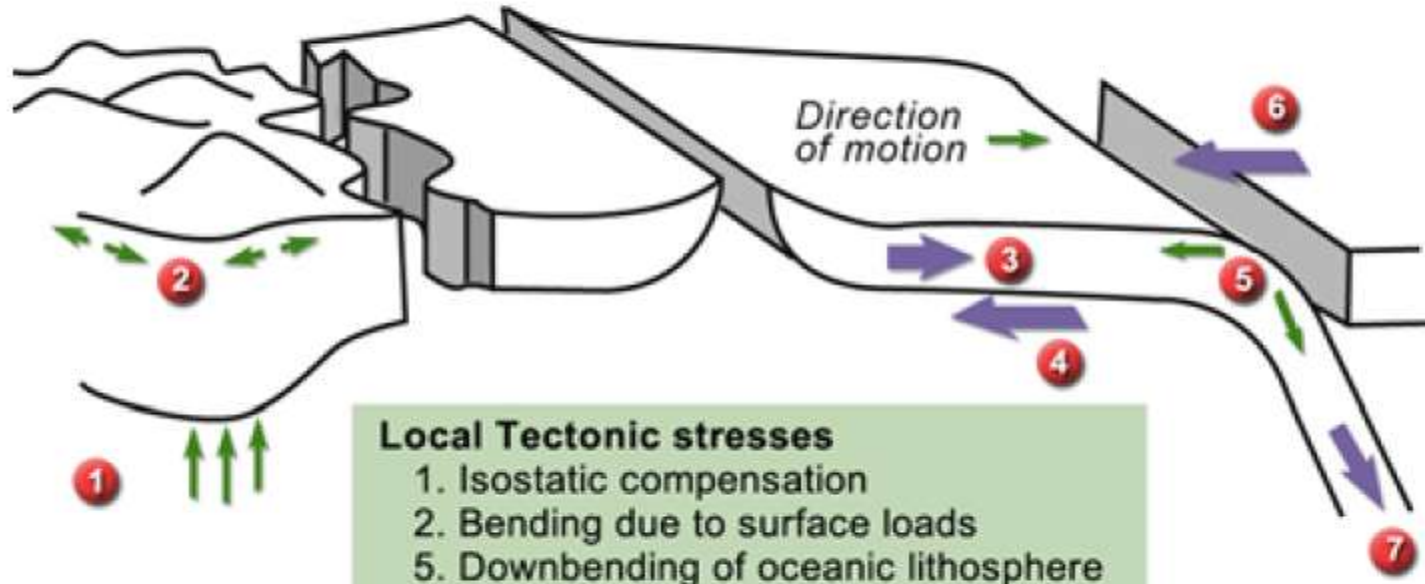


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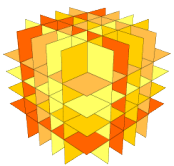


Sources of Tectonic Stress

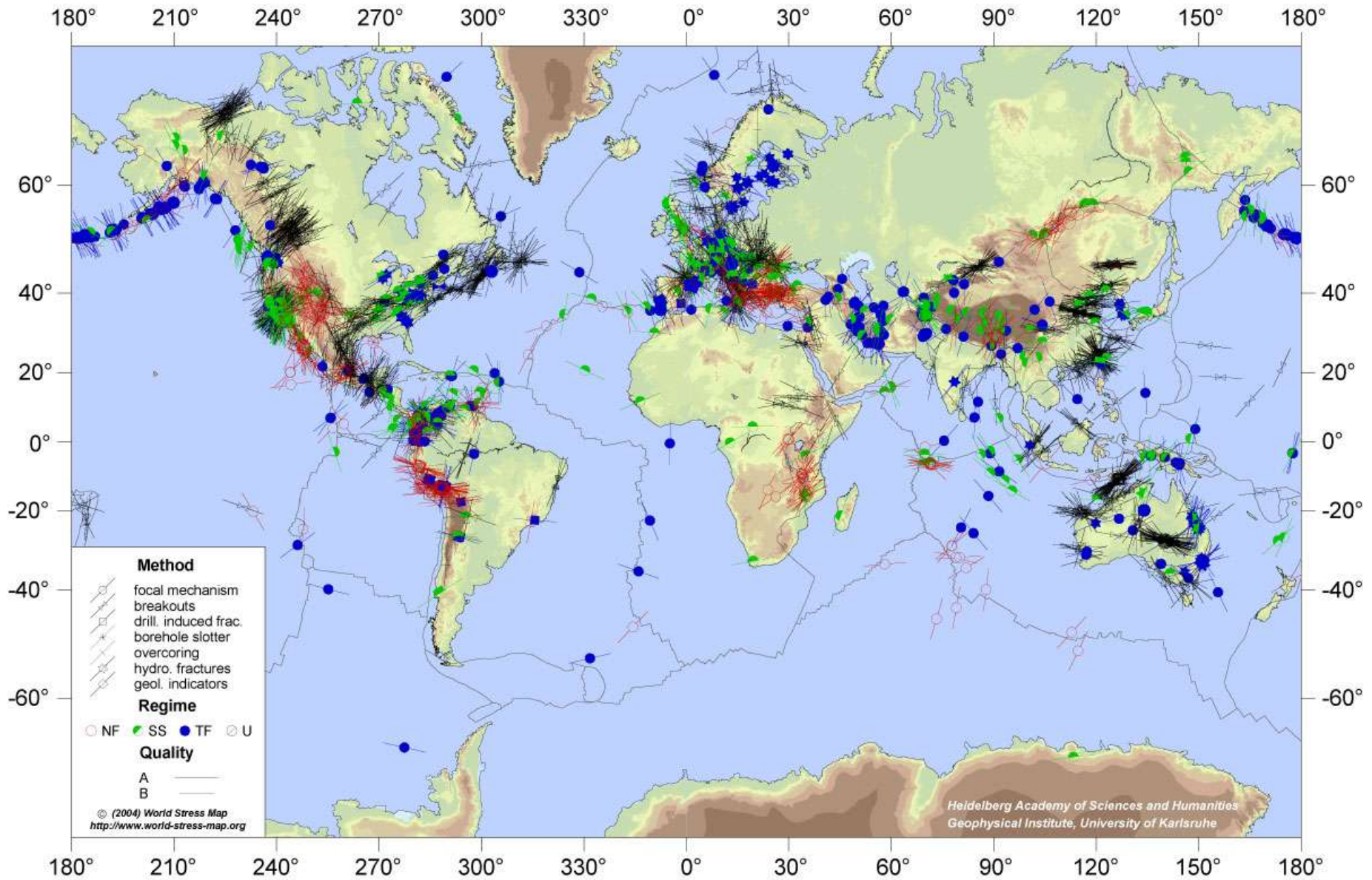


- Local Tectonic stresses**
- 1. Isostatic compensation
 - 2. Bending due to surface loads
 - 5. Downbending of oceanic lithosphere

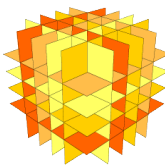
- Broad-scale tectonic forces**
- 3. Ridge push from oceanic ridges
 - 4. Shear traction at base of lithosphere
 - 6. trench suction on over-riding plate
 - 7. Net slab pull at subduction zones



World stress map

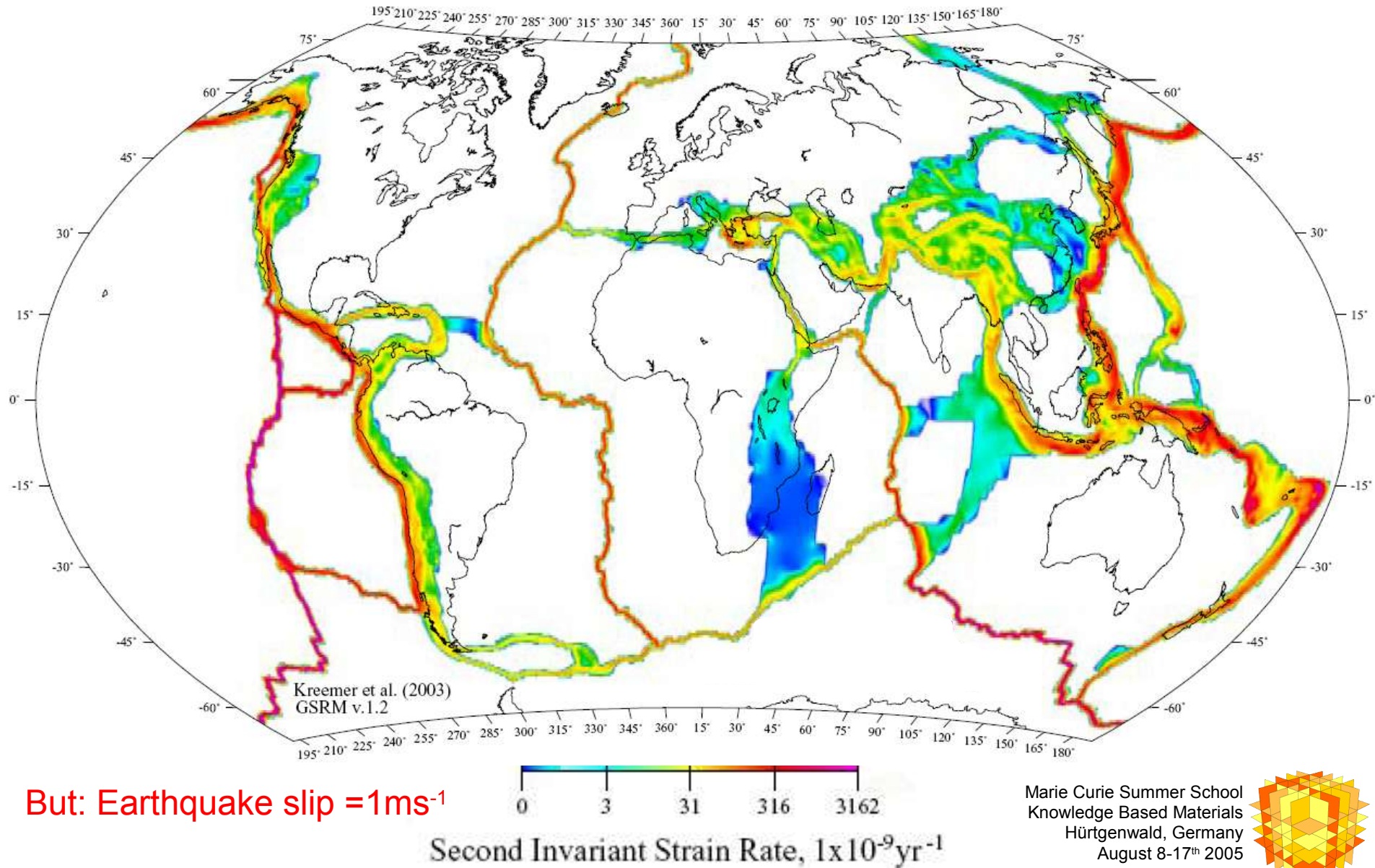


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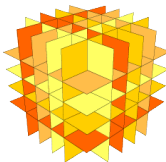


Time scales

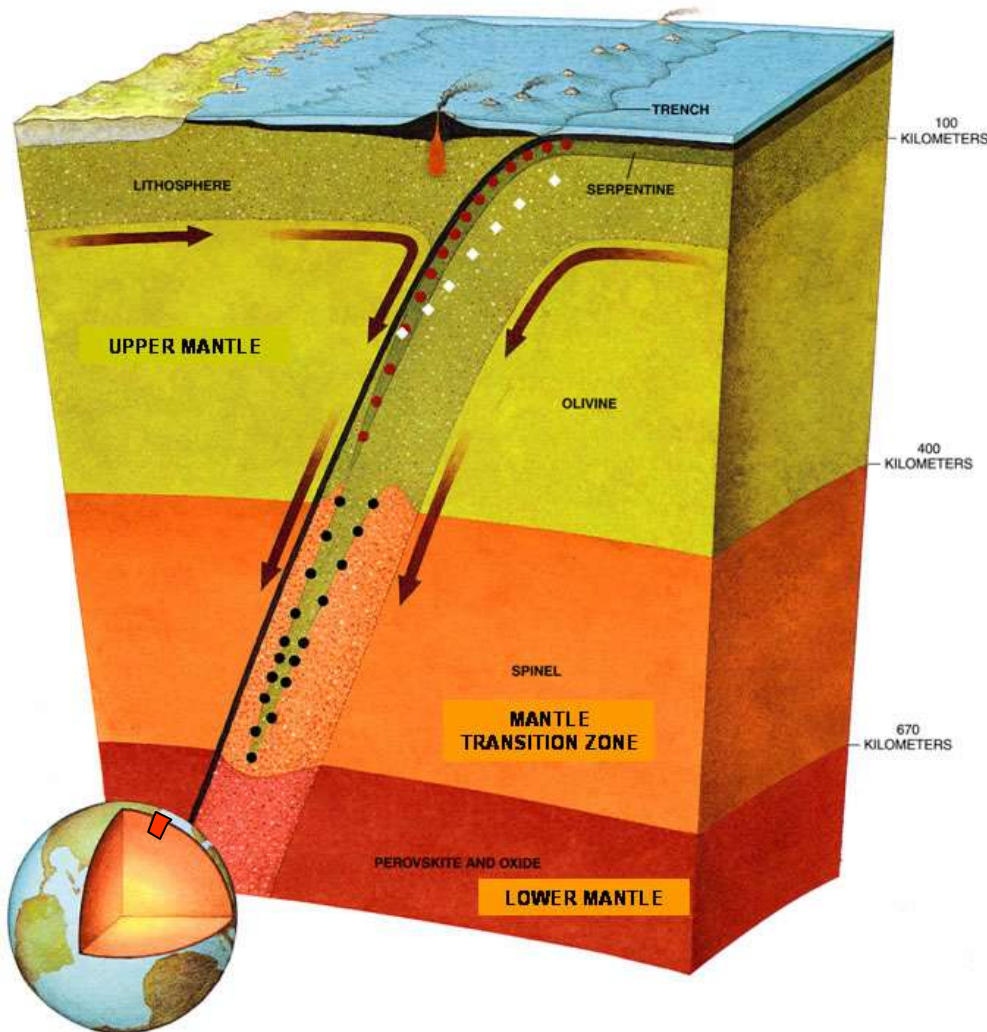
Absolute $10^{-9} \Rightarrow 10^{17} \text{s}$
Strain rates $10^0 \Rightarrow 10^{-15} \text{s}^{-1}$



But: Earthquake slip = 1ms^{-1}

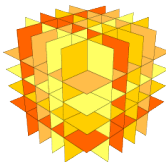


Length scales



Location	Depth km	Lithostatic Pressure
ocean ridge crest	2.5	25 MPa
abyssal plain	4	40 MPa
deep sea trench	12	120 MPa
base of crust	30	1 GPa
transition zone	600	20 GPa
Core-mantle boundary	2900	140 GPa
Center of Earth	6400	360 GPa

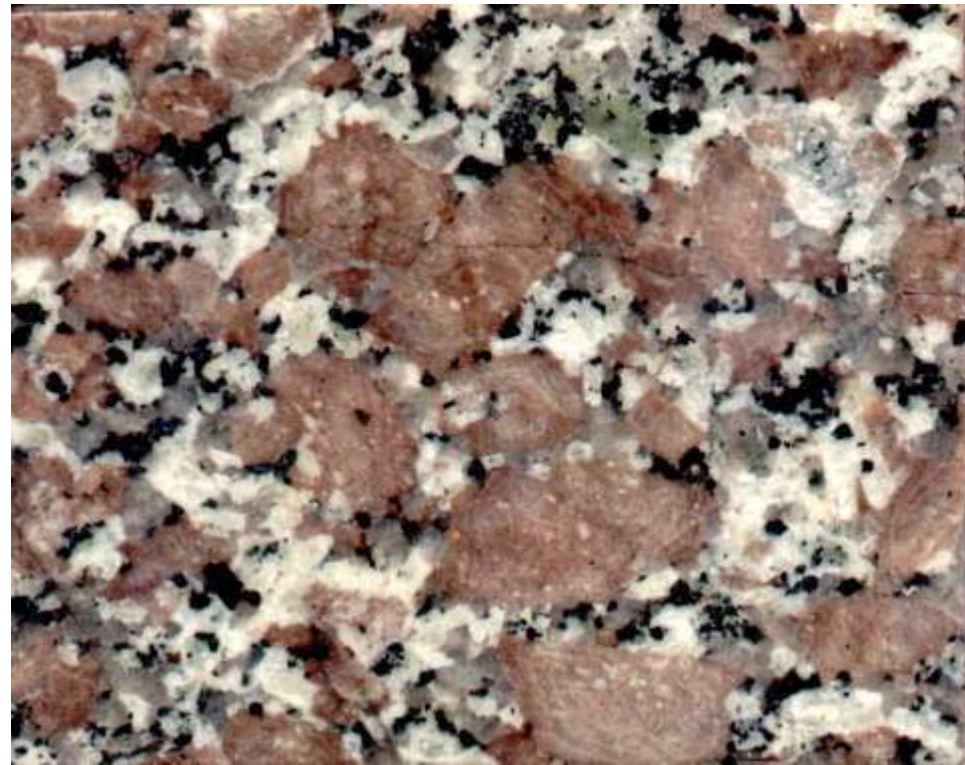
- Range $10^{-9} \Rightarrow 10^7$ m
- Crustal T gradient $0.02 \text{ } ^\circ\text{C m}^{-1}$
- Pressure Gradient
- Rough concentric compositional layering
- Experimentally, 25 GPa possible (1mm³ sample)
- Tectonic stresses probably of the order of 10-100 MPa



Earth Materials

Crust & upper mantle:

Element	Weight %
O	44
Mg	22
Si	21
Fe	6
Ca	2.5
Al	2.3
Na	0.3



- Mostly Silicates:

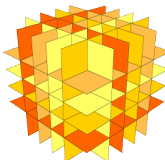
- >90 % Crust, >98% upper mantle
- KAlSi_3O_8 Monoclinic
- SiO_2 (Hexagonal, Trigonal)
- $(\text{Mg,Fe})_2\text{SiO}_4$ Orthorhombic

- Low symmetry \Rightarrow high elastic, plastic, surface energy, electrical anisotropies

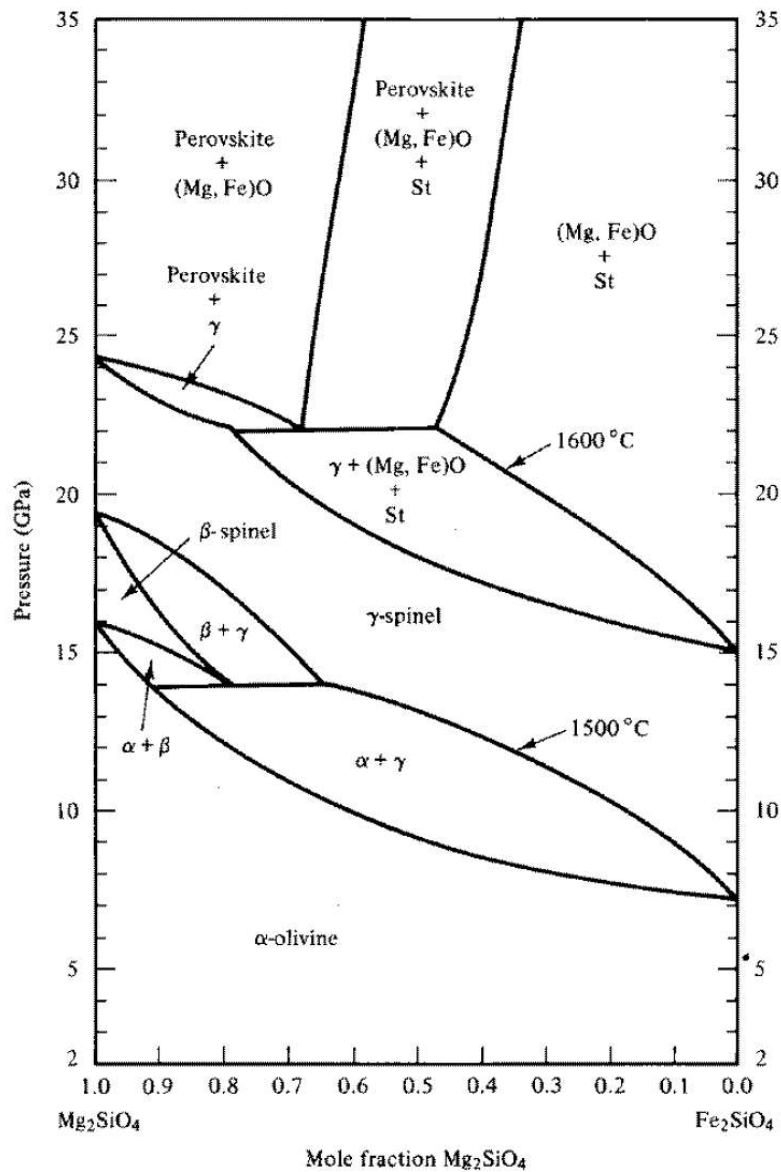
- Most rocks polyphase

- Fluid phases common (water, melt, CO_2)

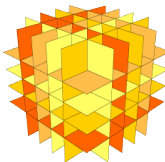
- Complex loading histories



Mantle Phases



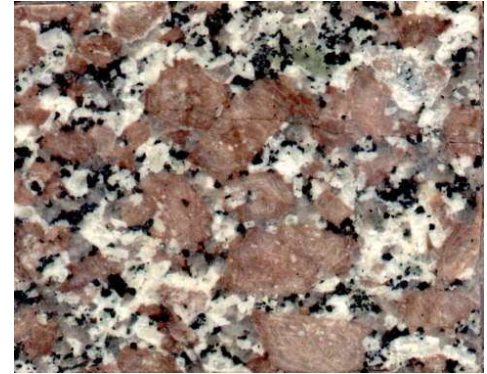
- Most of the Earth is dynamically recrystallising as we speak
- Phase changes important



How do rocks form:

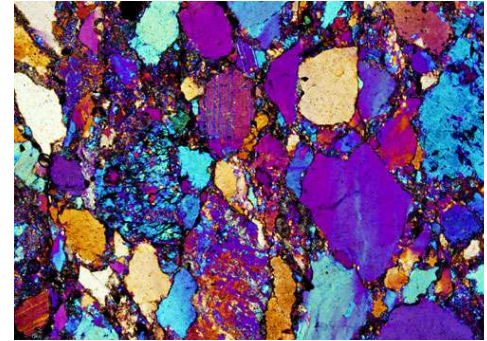
Igneous (solidifies from melts)

- Intrusive (Solidifies within an existing rock mass)
- Extrusive (Solidifies on Earth's surface)



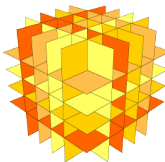
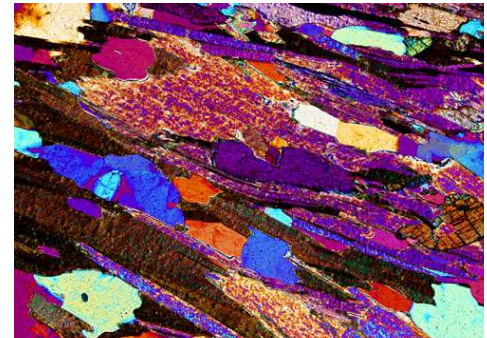
Sedimentary (destruction, transport & deposition)

- Mechanical
- Chemical
- Biological

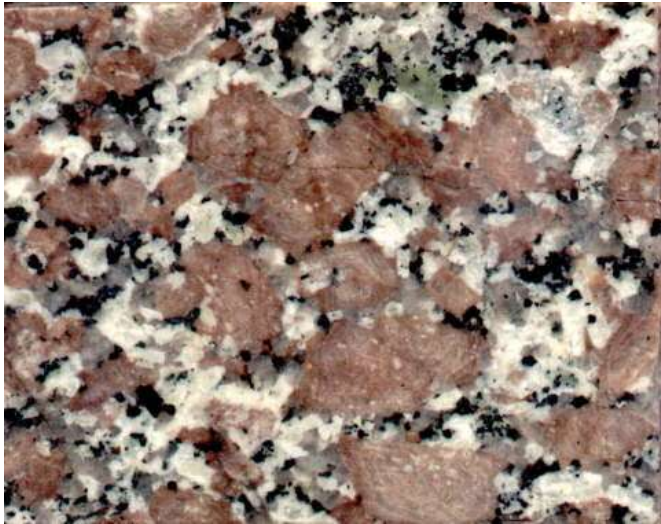


Metamorphism & Diagenesis

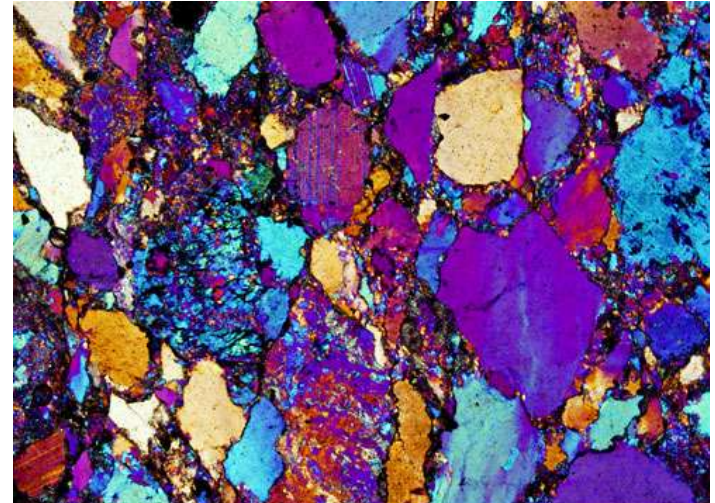
- Solid state (Nucleation, Chemical reaction, Phase change)
- Partial melt extraction
- Fluids (H_2O , CO_2)



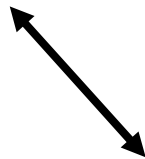
Cycle of rock formation



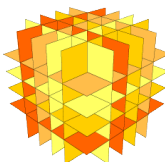
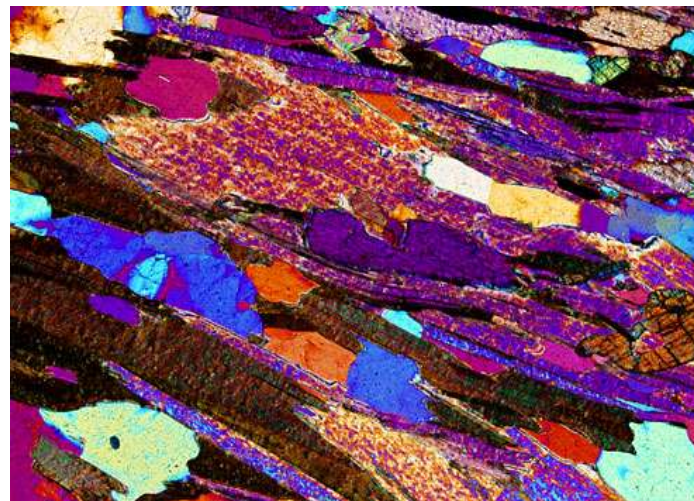
Igneous



Sedimentary



Metamorphic



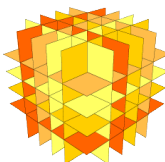
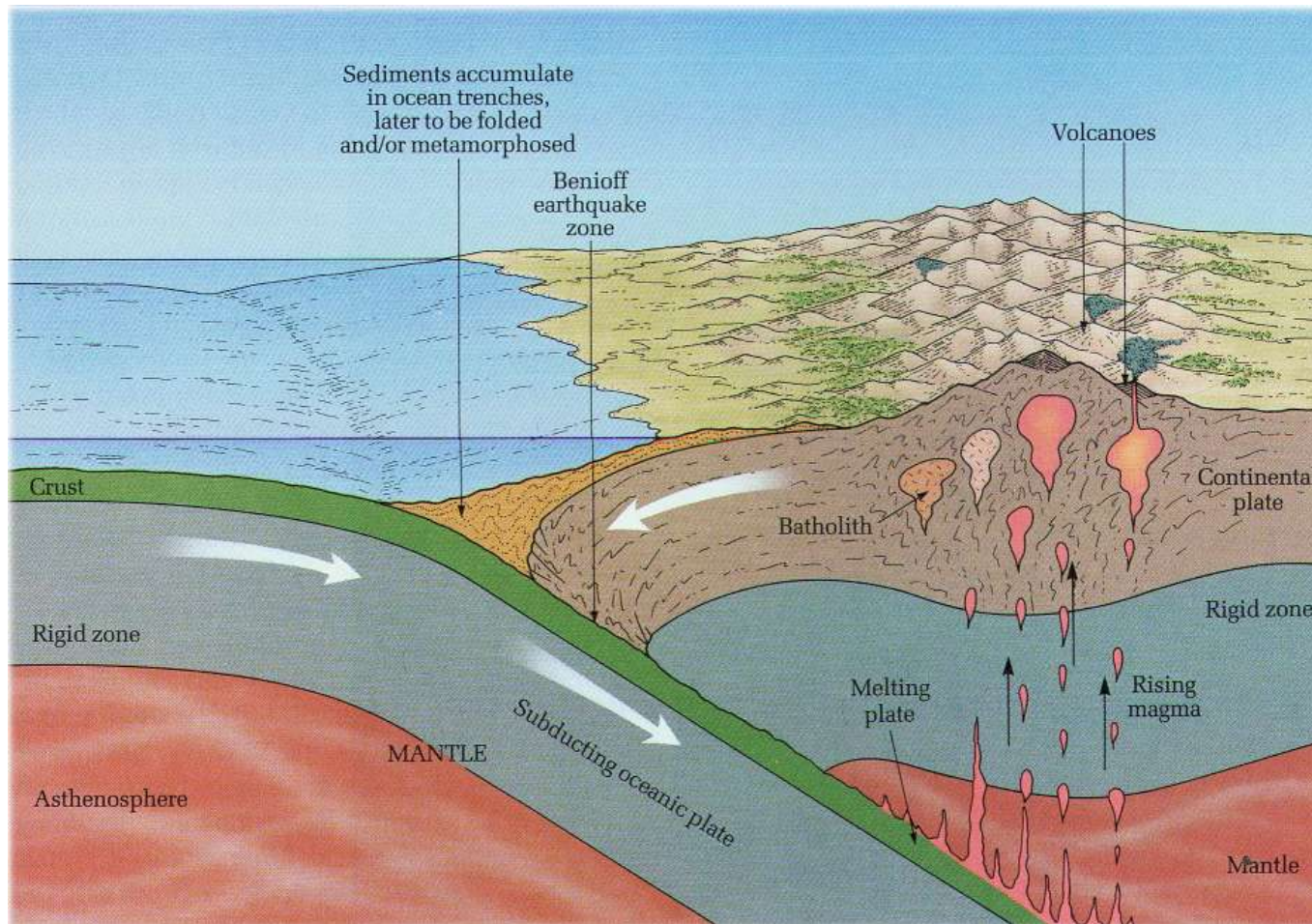
How do single phase rocks (s.l.) form?

Solidification from melt (Na,Ca Silicate)

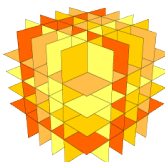
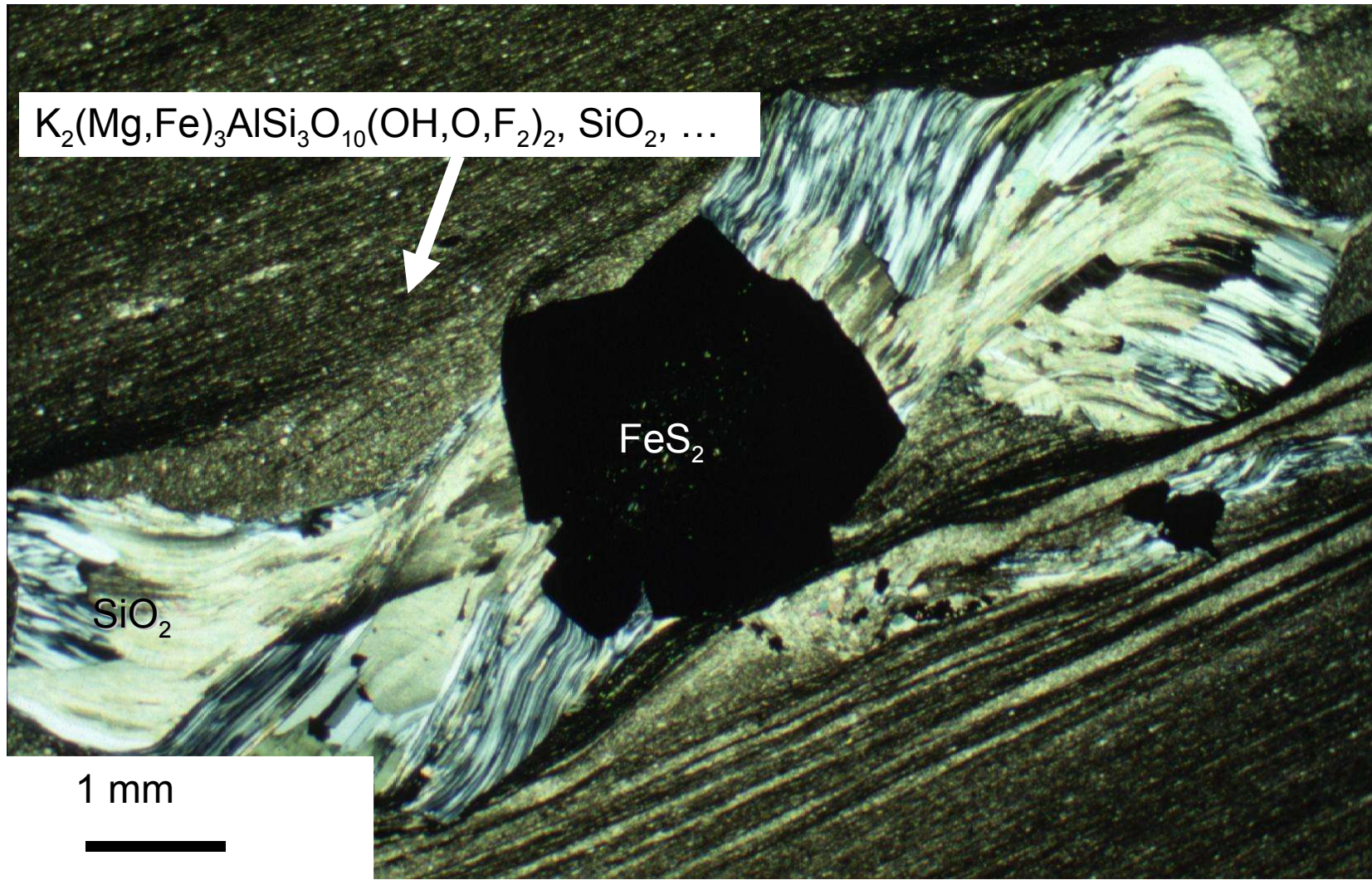
Fluid-Mechanical segregation (SiO_2)

Chemical segregation (CaCO_3 , NaCl, SiO_2)

Bio-deposition (CaCO_3 , SiO_2)



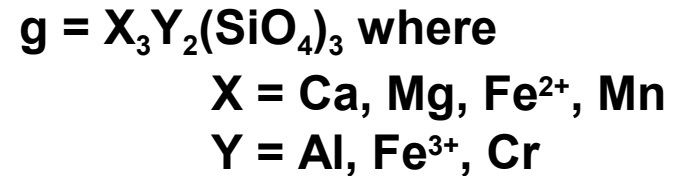
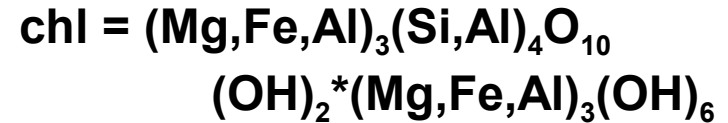
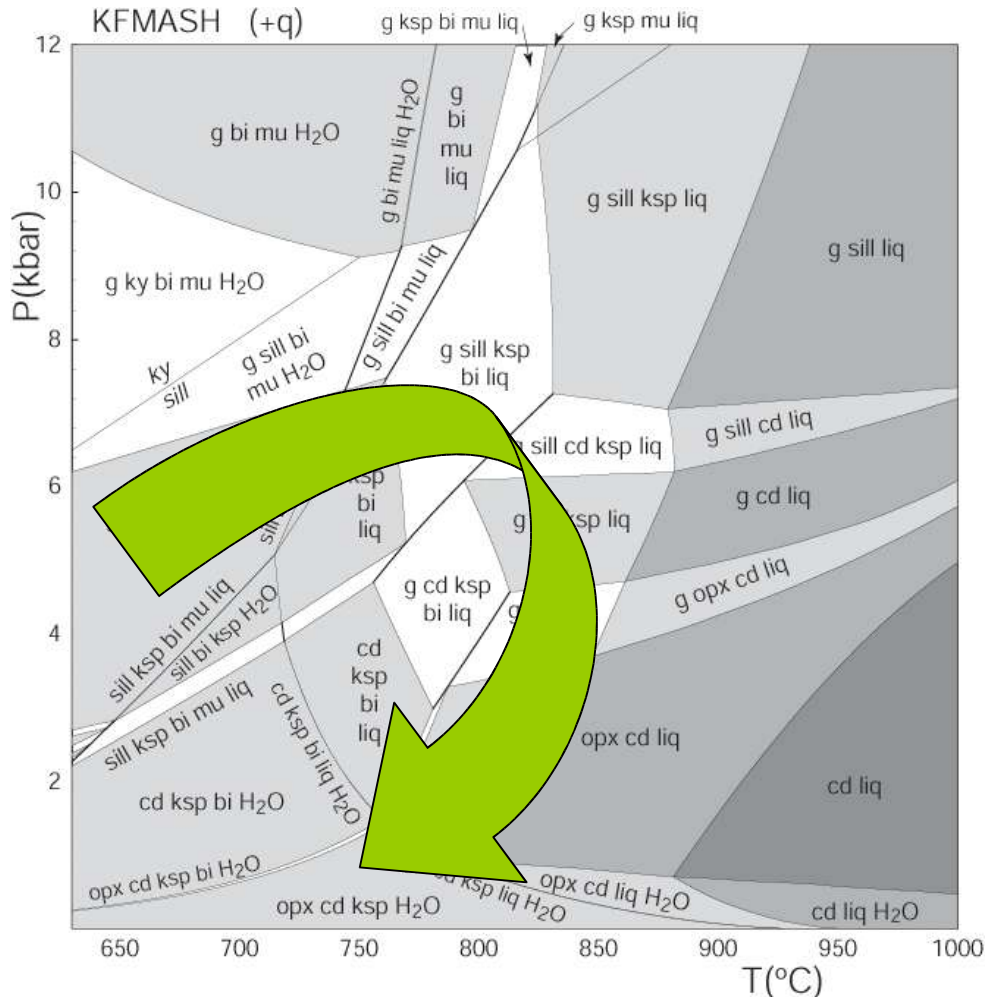
Chemical Segregation



Solid-state long-range reactions

chl-chlorite; bi-biotite; g-garnet; ky-kyanite; and-andalusite; sill-silliminite; st-stauralite; q-quartz

liq-liquid; opx-orthopyroxene; cd-cordierite; ksp-Potassium feldspar; mu-muscovite



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-
-

Powell 1997, typical mudstone mineral stability fields

